clay from the mainland. Organic matter is diluted by this detrital sediment so that it forms a lower percentage of total sediment in nearshore than in offshore basins. Even though hydrocarbons and porphyrins are also diluted by detrital sediments in nearshore basins, both are much more abundant (constituting a higher percentage of total sediment) in nearshore than in offshore basins. Thus, it is evident that both hydrocarbons and porphyrins are more easily oxidized than is total organic matter and that their preservation is greatly enhanced by rapid burial which removes them from contact with the oxidizing overlying water. Comparison with estimated ultimate oil recoveries from Los Angeles basin shows that far more organic matter, hydrocarbons, and porphyrins were produced and deposited in the basin sediments than were required to form the petroleum. Nevertheless, present production of petroleum from Los Angeles basin is at a rate that appears to be about 150,000 times greater than its rate of formation.

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Sources and Modes of Genesis of Nitrogen and Sulphur Compounds in Crude Oils

Nitrogen and sulphur compounds are present in all crude oils and, in some instances, may represent the major portion of the crude. While many sulphur compounds in crude oils have been identified, or the structure of the sulphur-containing group determined, relatively little is known concerning the structures of the nitrogen-containing compounds. Organic nitrogen is present in abundance in plant and animal detritus accumulating in aquatic sediments. From current knowledge of the structures of nitrogen compounds formed by plants and animals and conditions for post-depositional degradation, predictions can be made concerning the type of nitrogen compounds likely to be present in crude oils. It is improbable, however, that the quantity and variety of sulphur compounds present in crude oils could have been derived entirely from organo-sulphur compounds contained in the living source material. While derivatives of sulphur compounds synthesized by plants and animals may be present in crude oils, a larger portion of the sulphur compounds appearing in the crude must be formed after incorporation of the organic source material into the bottom sediments. Mechanisms proposed for the geochemical synthesis comprise reactions of plant—or animal-produced unsaturates and oxygen-nitrogen heterocyclics with sulphur or hydrogen sulphide formed by bacterial reduction of sulphonate ion.

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Bar-Finger Sands of Mississippi Delta

Elongate, lenticular sand bodies, termed bar fingers, characterize the Mississippi birdfoot delta sedimentary complex. They underlie the 15-20-mile-long principal distributaries or passes of the river which radiate from Head of Passes, and have formed in response to long continued distributary-mouth-bar deposition. These sand bodies attain a thickness of more than 250 feet and a width of as much as 5 miles. Their thickness results in part from subsidence brought about by compaction of underlying clays, and the width beneath a given pass is comparable to that of the actively forming distributary-mouth bar. The sand bodies comprise beds of well sorted, fine to very fine sand or silt and occasional thin layers of clay and silty clay. Diagnostic features are numerous cross-bedded thin layers in which the principal elements dip seaward, laminae composed of root and wood fragments, sand-size grains of lignite, and an absence of faunas. The bar-finger sands grade downward into delta-front silty clays which rest upon deeper marine prodelta clay deposits. Laterally they inter-finger with extensive thick clay sections which accumulated in delta-front, bay, and marsh environments. The sands are transitional with overlying natural levee, marsh, and bay deposits. Locally, the bar fingers have been disturbed by the upward movement of clays from underlying deposits to form mud lumps.

Bar fingers are distinguished from other sand bodies in the Recent deposits of the Mississippi deltaic plain by their greater thickness and their pattern of distribution. Their ancient counterparts have been recognized in the Pennsylvania Booch sands of Oklahoma. Abnormally thick sand masses present locally within the Eocene Wilcox and Sparta formations of Louisiana may also be deltaic bar-finger sands.

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A.A.P.G. Basement Project

The Basement Rock project of the A.A.P.G. was initiated in the fall of 1956; it is supported from the Research Fund. The Basement Project Committee, organized on a geographic basis, is currently compiling a basement map of North America between 24° and 60° N. Latitudes. This map will consist of two parts: (1) a map showing basement wells with code number, outcrops of basement rocks differentiated as to age and gross lithology, and contours on the basement surface and (2) a geologic and structural map of the basement. Preliminary copy for (1) is nearly complete. This map will be published through the cooperation of the U. S. Geological Survey and will be accompanied by a text giving basic data for all wells.

Basement studies are important in regional evaluations. Knowledge of basement geology and structural grain aids in interpretation of geophysical data. Movements along basement structures produce structures in younger basin rocks which can be prospectively more effectively if the basement control is recognized. Basement topography controls the facies of overlying sedimentary rocks. A regional knowledge of basement terranes is valuable in determining source of sediments and direction of transport. In areas where basement rocks are not "granite," thousands of unnecessary feet of hole have been drilled into metasedimentary and volcanic rocks. In some areas fractured basement rocks are reservoirs.

Petrographic methods, supported by geophysical and other information, can, within the limits of well control, establish (a) major lithologic and tectonic features such as orogenic belts, volcanic terranes, plutonic terranes, and fault zones and (b) tectonic divisions within concealed orogenic belts, such as allochthonous plates, belts characterized by different type and degree of metamorphism, and zones of igneous activity.

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Oil Exploration in Green County, Kentucky, and Adjacent Areas
New oil discoveries and expansion of Green County oil field have made 1959 the best year in the history of oil production in Kentucky with a total of 21,007,141 barrels produced to end of September. Green and Taylor counties had produced respectively 8,015,382 and 548,462 barrels in the same period from approximately 2,000 wells at depths from 350 to 500 feet. The Silurian Laurel dolomite, the main pay of the field, is found at depths of a few feet to about 60 feet below the New Albany black shale. Intracrystalline and vuggy porosity averages about 12 per cent and ranges up to 16 per cent. Permeabilities vary from a few up to 2,500 millidarcys. Thickness of the pay varies from a few feet to 25 feet, with 12 feet common.

An abundance of highly saline water produced with the oil from the field has posed a serious pollution problem of fresh-water supplies. The Kentucky Water Pollution Control Commission is making progress toward correction of the situation.

The search for oil has spread throughout central and eastern Kentucky on both flanks and crest of the Cincinnati arch. New oil fields and pool extensions have been found in Lincoln, Metcalfe, Allen, Barren, Hardin, Simpson, Cumberland, and Clinton counties. These relatively shallow areas produce from Lower Mississippian, Devonian, Silurian, and Ordovician rocks. Porosity zones truncated by New Albany black shale and fracture porosity in the Ordovician rocks form most of the oil traps. Some reef production is present in the Granville pay of the Ordovician in Clinton County.

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Geologic Mapping of Submerged Continental Margins

In the late nineteenth century it was discovered that rocks dredged by fishermen from the continental slope contained Cretaceous and Tertiary fossils. In the mid-1930s, Stetson systematically dredged such rocks from the Georges Banks' submarine canyons, concluding that the canyons had been cut deeply into a presumed sequence of "foreset" and "topset" beds.

Subsequently ancient rocks have been discovered on linear segments of the continental slope not near submarine canyons and it now appears that the continental slope from Newfoundland to Puerto Rico forms a continuous outcrop of Tertiary and Cretaceous sediments.

Echo-sounding profiles of the continental slope show a succession of topographic benches and gradient changes which have been correlated in a few areas with the outcrop pattern determined either by coring and dredging or by extrapolation from nearby drill holes. Scattered and less complete data from other parts of the world suggest that continental-slope outcrop benches are of common, if not of general occurrence. By correlating benches dated and verified by dredging, it is possible to construct geological maps of the continental slopes. In addition to greatly adding to the geological understanding of the continental shelves, extensive continental slope outcrops are of great significance in such major geological problems as the origin of the continental slopes and the origin of the continents and oceans.

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Scientific Objectives of Mohole, and Predicted Section

The major scientific objective of the AMSOC project is to obtain samples of the upper part of the earth's mantle and to determine the nature of the Moho discontinuity. Boisse's analogy of 1850 suggesting that meteorites were a fair approximation of the composition of the earth's interior was a brilliant idea for its day, but is too inexact for present purposes.

Above the discontinuity is a layer commonly called the "crust" and generally considered to be basalt. It is extraordinarily uniform in thickness suggesting that its base represents a phase transition and that the Moho under the oceans might represent the level of some isotherm, or past isotherm, at which a reaction took place.

Above the "crust" is a layer of variable thickness and seismic velocity thought to be consolidated sedimentary or volcanic rocks. Finally one comes to the unconsolidated sediments of the ocean floor which are a few hundred meters thick and no doubt resemble the material obtained from shallow cores. In these last two layers one might hope to find fossils going far back in the history of the oceans and to derive from this record information of extraordinary scientific importance.

The following predictions are made.

(1) The mantle will be peridotite resembling the olivine nodules found in basaltic volcanoes and St. Paul's rock.

(2) The "basalt crust" will be serpentinitized peridotite, hydrated mantle material.

(3) The Moho discontinuity represents a past isotherm above which serpentine was a stable phase.

(4) The sedimentary column will be very incomplete and have many great hiatuses.

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Geometry of Producing Mesaverde Sandstones, San Juan Basin

Within the San Juan basin the sandstone zones that occur at the top and bottom of the Mesaverde group were not deposited as a continuous blanket sand. In some areas thick, relatively clean sandstone units occur. In other areas thin, poorly sorted sandstone zones are found. These sandstone units exhibit a definite geometric pattern of distribution. Sandstone beds of the Point Lookout formation were deposited as a shoreline phase of a sea regressing northeastward. Sandstone zones of the Cliff House formation represent the shoreline of a sea transgressing southward at a later date. The shoreline along which these sands were deposited moved rapidly across some areas. In other areas, it remained stationary for relatively long periods of time. The thicker sands correspond to places where the shoreline remained stationary for the longer periods of time.

The successive positions of the various Cliff House and Point Lookout shorelines have been established both vertically on cross sections and laterally on maps. Those positions where the shoreline stabilized for relatively long periods of time are apparent in the form of "steps" that can be traced across the central part of the San Juan basin. The relatively thick, well sorted sandstone units that correspond with the positions where the shoreline stabilized have been divided into a series of "benches" of varying widths. Excellent examples of major "steps" in the Cliff House shoreline can be seen in surface exposures in the southeast and northwest parts of the San Juan basin. These "steps" exposed at the surface in the northeast part of the basin exhibit the same strandline trend and in general are correlative with the "steps" found in the subsurface.