

**A mineralogical and fluid inclusion study of the West Gore Sb-Au deposit, central Meguma Terrane, Nova Scotia:  
a possible high-level analogue for mesothermal Meguma Group lode-gold deposits**

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Mineralization at the West Gore Sb-Au deposit (7158 oz. Au produced), hosted by NE-trending slates of the Halifax Formation, occurs within three, subparallel and subvertical, NW-trending fault structures. Vein fill is dominated by unstrained, bimodal quartz (coarse prismatic and fine-grained equant) and paragenetically later carbonate+chlorite and associated ore minerals occupying vugs or orthogonal fractures. The ore assemblage consists of early stibnite, locally replacing po, with associated po  $\pm$  py  $\pm$  arspy  $\pm$  Fe-rich sph  $\pm$  cpy  $\pm$  native antimony. Stibnite occurs either as coarse blebs or delicate, acicular to reticulate grains intergrown with quartz and/or carbonate. Stibnite is replaced by a complex, concentrically zoned intergrowth of native Au (<1 wt. % Ag) and colloform-textured Au-Cu alloy; this assemblage is frequently accompanied by zones of carbonate  $\pm$  chlorite enrichment. The latest stage of mineralization consists of Fe- and Mn-oxides and secondary Sb minerals [kermesite ( $\text{Sb}_2\text{S}_2\text{O}$ ) and valentinite ( $\text{Sb}_2\text{O}_3$ )].

Fluid inclusions (quartz only; 5-15 $\mu$  size) occur as (1) isolated groups or (2) planar arrays along healed microfractures. The cores of prismatic-quartz grains are generally inundated with inclusions whereas the margins are inclusion-free; hence, a probable primary or pseudo-secondary origin for the inclusions. Inclusion types are: (i) V-rich ( $\text{H}_2\text{O}$ ), (ii) L-rich ( $\text{H}_2\text{O}$ ), and (iii)  $\text{H}_2\text{O}-\text{CO}_2$  ( $X_{\text{CO}_2}=0.1$ ). While type (iii) is rare, (i) and (ii) are abundant and their mutual coexistence suggests fluid unmixing. Thermometric data ( $^{\circ}\text{C}$ ) are summarized as follows: (1)  $T_m(\text{CO}_2)=-58.7\pm 1.0$  (n=8); (2)  $T_m(\text{ice})=-4.0\pm 1.0$  (n=50; ca. 5-8 wt. % equiv. NaCl); (3)  $T_b(\text{CO}_2)=25.4\pm 0.8$  (n=4; to L) and 13.3 (n=1; to V); and (4)  $T_b(\text{H}_2\text{O})=\text{ca. } 200-400$  (n=177; to L) and 370 (n=5; to V). In addition, (1) clathrate formation indicates ca. 1 mol % dissolved  $\text{CO}_2$  in type (ii) inclusions, (2) decrepitation is common between 300-400 $^{\circ}\text{C}$ , (3) there is no correlation between salinity and  $T_b(\text{H}_2\text{O})$  and (4) ambient pressure of formation is estimated at 0.6 kbars (ca. 1700 m lithostatic cover).