

Both sandstone bodies are elongate, channel-fill, belt-type deposits of fluvial origin. The Hockingport Sandstone is a north-south oriented deposit up to 90 feet thick, blanketing an area of 225 square miles. The longer axis of the Waynesburg Sandstone is oriented N 20° E, the maximum thickness is 75 feet, and the preserved deposit blankets an area of 880 square miles. The grand mean of dip directions of cross-stratification readings of the Hockingport Sandstone is N 19° W and of the Waynesburg Sandstone, N 10° E.

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UPHEAVAL DOME, SAN JUAN COUNTY, UTAH

Upheaval Dome is a dissected domal structure that lies in the rugged "Canyon Lands" of southeastern Utah. Extensive canyon development within the dome's immediate area has produced remarkable exposures of the feature. It is surrounded by a distinct rim syncline and piercement is much in evidence within the dome's central portion; the area of deformation is approximately three miles in diameter and a vertical displacement of at least 1,200 feet exists at the feature's center. Strata that crop out within the domal area range in age from Permian to Jurassic, and all exhibit deformation by the forces that produced the dome. The area is underlain by some 3,000 feet of Pennsylvanian saline strata. The origin of the dome has not been proven, but the following hypotheses have been advanced: (1) it is a cryptovolcanic feature; (2) it was formed by meteorite impact; and (3) it is a salt dome (but there are conflicting views concerning the reason for salt intrusion at this specific site). An additional hypothesis for salt intrusion is advanced: Upheaval Dome is the product of salt intrusion resulting from differential compaction of the sediments lying above and on the flanks of a low conical hill on the Precambrian surface.

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CRETACEOUS STRIKE-VALLEY SANDSTONES, NORTH-
WESTERN NEW MEXICO

Cretaceous sandstone bodies, which are the principal oil reservoirs in the San Juan Basin, have been interpreted as offshore bars related to deposition of the regressive Gallup Sandstone. A detailed outcrop and subsurface study in the northwestern San Juan Basin indicates that these sandstone bodies are strike-valley sandstones in the transgressive marine "basal Niobrara" unit which rests unconformably on the Gallup Sandstone and older units.

Paleogeography of the pre-Niobrara unconformity consisted of northwest-southeast trending cuestas and intervening strike valleys with local relief of over 100 feet. Cuestas and valleys are related to the subcrop of alternating resistant and non-resistant units in the truncated sequence. Elongate, lenticular sandstone bodies overlying the unconformity occur in strike valleys on the northeast side of cuesta scarps. Individual sandstone bodies, with a maximum thickness of about 50 feet, lap out to the southwest against the cuestas and thin to the northeast by facies change to shale. Younger sandstone bodies extend progressively farther to the southwest.

Basal Niobrara sandstones are fine to coarse grained and glauconitic, and contain marine microfossils. Interbedded and laterally equivalent shales also contain marine fossils, including both benthonic and pelagic forms. The sandstones are characterized by broadly

lenticular sets (up to 6 feet thick) of high-angle cross-stratification. Measurements of cross-stratification dip directions at 20 localities indicate transport by currents flowing generally southeastward.

These sandstones are best interpreted as nearshore marine sands deposited in strike valleys on the seaward side of cuesta scarps during a general transgression to the southwest. During the transgression, cuesta ridges acted to stabilize temporarily the position of the shoreline, permitting accumulation of sand nearshore while clays were deposited farther offshore to the northeast. Transport of sand was largely in the form of underwater dunes migrating alongshore to the southeast.

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AN APPLICATION OF ELECTRONIC DATA PROCESSING
TECHNIQUES TO PALEONTOLOGY AND STRATIGRAPHY

A method for electronic data processing of paleontological/stratigraphic information has been devised and is in operation. Through use of a species accession reference number, paleontologic, stratigraphic, and ecological data (including synonyms of species) can be collated to reveal essential information on a given form. In addition, retrieval can be effected through use of the genus and species names as a retrieval code, although this is a slower process.

A bibliographic accession reference number allows complete treatments of bibliographic data, including abstracts, faunal lists, and other important collateral information. This bibliographic accession reference number is added to the species index cards to effect a cross-reference and corollary check of data.

This system is in operation for the McLean Card Catalogue of American Foraminifera, the Card Catalogue of Ostracoda, and the H. S. Puri Card Catalogue of Recent Ostracoda, all of which are completely cross-indexed by the system. With minor modifications, the index can be used for museum types and collections; it will be used for numerous other applications. The basic file will be available, through mechanical reproduction, to anyone desiring to use it.

Electronic data processing methods will partially replace, and greatly augment, manual methods; but for most operations, manual searches will still be the most efficient method of retrieval. Computers are advised only for extensive searches and, periodically, to correct and arrange the files. The interpreted, mechanically reproduced file cards are efficient for manual use.

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GENERAL GEOLOGY AND HYDROCARBONS OF THE
NORTHERN AMADEUS BASIN, AUSTRALIA

The Amadeus basin is a structural depression extending easterly from the Canning basin in Western Australia to the Great Artesian basin in Queensland. It covers about 80,000 square miles and contains up to 30,000 feet of late Proterozoic and early Paleozoic marine and continental sediments.

The marine cycle of deposition started in late Proterozoic time and terminated in late Ordovician. No unconformity is present at the base of the Paleozoic sequence and both Proterozoic and Paleozoic beds are unaltered except along the northern margin of the basin where low grade metamorphic facies were developed in