

considered in hypotheses that explain the development of continental margins.

WARREN J. WAHLSTEDT, Cities Service Oil Co., Bartlesville, Okla.

#### ECONOMICS, DECISION MAKING, AND EXPLORATION POLICY

The objective of exploration is to find petroleum and increase the company's reserves. To accomplish this goal, the company must utilize its resources (*i.e.*, money and personnel) in the most effective way possible. This might seem to indicate that, to accomplish this goal, exploration should be restricted to those areas in which the potential big fields (*e.g.*, offshore) are being discovered. This is incorrect, however, because a company has limited resources and may not be able to obtain favorable leases in such areas. If this is the case, any effort spent on these areas will be a fruitless dissipation of resources. This does not mean that a company should not compete in these areas, but rather that it should learn how and where it can compete effectively.

Independents do not face this problem because their resources are so limited they cannot have illusions about how and where to compete. Very large companies can overcome this difficulty by overbidding on areas of interest, thus insuring that they can obtain their goals. It is the companies which fall between these extremes which face the problems of how and where to compete.

In attempting to compete, most companies use the trial-and-error method and in the process dissipate resources which they cannot afford to expend. Another alternative is to use the computer. Algorithms have been developed to simulate decision making in conditions of uncertainty; some have been done concerning exploration and lease bidding. By use of this type of approach and adapting the OR method to the problems of exploration, a company can find how and where best to invest its resources.

ROGER G. WALKER, Dept. Geology, McMaster Univ., Hamilton, Ont.

#### RIPPLE-DRIFT CROSS-LAMINATION IN TURBIDITES

Ripple-drift cross-lamination is the name given to a series of ripples which "climb upon the backs" of each other as a result of the addition of sediment from suspension during ripple migration. The type found in turbidites is characterized by continuity of lamination across the ripple system, changing composition of laminae from lee side (mud) to stoss side (silt and sand), and gradual upward decrease of ripple amplitude.

Geometric analysis has shown that the angle at which the ripples climb is a function of lamina thicknesses on the lee and stoss sides, the angle of the lee and stoss slopes, and the symmetry of the ripples. Computation of many angles of climb and different ripple geometries has shown that all three factors (thickness, angle, and symmetry) are equally important in determining the angle of climb. Field measurements of ripple-drift morphology in Ordovician turbidites of the Gaspé Peninsula, Quebec, indicate angles of climb from about 3 to 40°, with cosets of ripple-drift cross-lamination ranging in thickness from 4 to 37 cm. There is a direct correlation between coset thickness and angle of climb, and lee-side lamina thicknesses tend to be two to four times greater than

stoss-side thicknesses. The factor controlling the angle of climb is the ratio of volume of sediment deposited from suspension to the volume of sediment moved on or very close to the bed and deposited on the lee sides. The velocity at which the ripples move downstream is also a function of the same ratio, plus a function of the hydraulic parameters controlling the formation of the ripples. Because of the complex interactions of these variables, it is not yet possible to estimate accurately the time taken for formation of the ripple-drift cosets. Crude estimates suggest a period of less than 1 or 2 hours for a coset 40 cm thick.

HAROLD R. WANLESS, Dept. Earth and Planetary Sciences, Johns Hopkins Univ., Baltimore, Md.

#### SEDIMENTARY STRUCTURE ZONATION ON TIDAL LEVEES, ANDROS ISLAND, BAHAMAS

The natural levees of tidal channels on northwestern Andros Island, Bahamas, show how primary sedimentary structures record small-scale variations in elevation and lateral position on a tidal flat. These levees, like stream levees, are highest adjacent to the channels, crest less than 50 cm above mean high water, and slope gradually for 25–300 m into shallow subtidal ponds.

The most pronounced variation in sedimentary structures is between the well-laminated sediments of the levees and the bioturbated sediments of the pond. In the intertidal and subtidal ponds, browsing and burrowing animals completely destroy primary lamination, but the long periods of desiccation on the levees keep burrowers out of the sediments.

Within the levees there are three distinct zones: (1) the levee crest, 30–50 cm above MHW, adjacent to the channel has smooth, parallel laminations 1–5 mm thick; mudcracks, where present, are discontinuous; (2) the central part of the levee, 10–30 cm above MHW, has thinner laminations, mostly less than 1 mm, that are disrupted by shallow mudcracks 2–6 cm apart. Cornflake-size chips and larger clasts are abundant; (3) on the pond side of the levees, 0–10 cm above MHW, thin, crinkled laminations less than 1 mm thick alternate with thicker laminations that have a prismatic structure inherited from surface mats of the blue-green alga *Scytonema* sp. Small mudcracks (2–8 mm) disrupt the lamination.

WILLIAM C. WARD, Dept. Geology, Rice Univ., Houston, Tex., ROBERT L. FOLK, Dept. Geology, Univ. Texas, Austin, Tex., and JAMES LEE WILSON, Dept. Geology, Rice Univ., Houston, Tex.

#### SURFICIAL ALTERATION OF PLEISTOCENE(?) LIMESTONE ADJACENT TO SALINE LAKES, ISLA MUJERES, QUINTANA RÓO, MEXICO

Two interesting rock types result from calichification and alteration of the weathered surface of bedrock limestone adjacent to shallow saline lakes on Isla Mujeres, Quintana Róo, Mexico.

Jet-black micritic limestone is produced on the periphery of the saline lakes, apparently as the combined result of calichification, blue-green algal penetrations, and sulfate-reducing bacterial action. Micrite layers are added by calichification. The black color is attributed to finely disseminated organic material and iron sulfide(?) produced by bacteria. Bacterial action not only may be partly responsible for the black color of the limestone, but also may account for the small amount