

CENE-OLIGOCENE, SANTA YNEZ MOUNTAINS, CALIFORNIA

Facies relations between deep and shallow-marine to continental deposits in the Eocene-Oligocene sequence of the Santa Ynez Mountains, California, have been studied in detail. The rocks studied include the Anita, Sierra Blanca, Juncal (with Camino Cielo Member), Matilija, Cozy Dell, and "Coldwater" formations. The topmost unit of the Eocene-Oligocene sequence, the nonmarine Sespe Formation, was not included in this study.

In landward sequence, the facies recognized include turbidites and marine lutites, proximal turbidites, shallow-marine, coastal, and continental facies. These are present in 2 major regressive sequences. In the first, the Juncal-Matilija sequence, thin-bedded turbidites and marine lutites are overlain by, and are laterally equivalent to, very thick proximal turbidites which pass upward into shallow-marine and coastal sands. The major sand accumulations are in the basin-margin shallow-marine, coastal, and proximal-turbidite facies. The second regression, the Cozy Dell-Sespe sequence, lacks significant proximal-turbidite deposits, but has extensive shallow-marine and coastal deposits. Facies distribution and stratigraphic sequence are explained as responses to the interplay of depositional and structural processes.

Detailed stratigraphic mapping has clarified correlation in the Eocene sequence of the Santa Ynez Mountains. Micropaleontology is used in support of correlations and bathymetric interpretations. Stratigraphic and paleontologic data resolve a biofacies problem in the lower to upper Narizian interval and clarify definition of the Eocene-Cretaceous boundary. Bathymetric interpretations based on comparison of fossil assemblages with modern Gulf of Mexico fauna are in better agreement with depths of deposition interpreted from lithologic data than interpretations based on comparison with modern assemblages in the Pacific Ocean off California.

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SEISMIC EXPLORATION IN CANADIAN ARCTIC

No abstract available.

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SEDIMENT DISTRIBUTION IN SOUTHWESTERN INDIAN OCEAN

Sediment distribution patterns in the southern Mozambique Channel and adjacent southwest Indian Ocean were investigated from short gravity cores and surface bottom samples. These shelves are covered by terrigenous sediments of varying grain sizes, chiefly silty, and upper slopes by hemipelagic calcareous mud. Areas of nondeposition on parts of the African shelf and upper slope reflect the winnowing action of the Mozambique-Agulhas Current. Relict foraminiferal faunas and exposed beachrock on the African shelf indicate a former lowered sea level. Lower continental slopes (below 1,500 m), continental rises, and plateaus are covered by foraminiferal marl and chalk oozes whose distribution correlates in part with the grain-size distribution. A size analysis of planktonic Foraminifera indicates that concentrations of small and large tests in some oozes are due to sorting. Manganese nodules are abundant on the Mozambique Plateau, which is swept by an eddy of Antarctic Intermediate Water. Natal basin sediments include turbidites, derived from neritic and bathyal depths, which are dispersed from northwest to southeast. Surficial turbidite layers did not reach the southern part of the basin floor, which is covered by pelagic clay and manganese nodules.

Average sedimentation rates of sediments younger than approximately 4,000-6,000 years range from 1 to 9 cm/1,000 years, depending on topographic position and distance from land. The average rate of accumulation for undisturbed deep-

sea ooze during the entire late Quaternary is 1.5-2.5 cm/1,000 years.

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TRANSPORT PROCESSES FOR LOWER PALEOZOIC RESEDIMENTED CONGLOMERATES OF APPALACHIANS

Three fining-upward sequences of resedimented conglomerates, totaling about 40 m in thickness, are present in the lower Paleozoic Cap Enragé Formation of Gaspé, Quebec. Boulders, dominantly of shelf-type carbonate rocks, are up to 3 m in diameter, but average about 10-50 cm. Individual beds are more commonly massive than normally graded, although inverse grading occurs at the base of many beds. Stratification, delineated by alternations of grain size, is common.

At the base of each fining-upward sequence, beds tend to be poorly sorted, and contain very large boulders. At the top of the sequences, well sorted conglomerates, with pebbles normally less than 5 cm, are interbedded with sandstones. Deep scouring is uncommon.

The best clue to the transport process which resedimented the conglomerates into the Appalachian geosyncline is the well-defined pebble fabric. Most beds show a strong preferred southwest orientation of *a*-axes in plan view. In vertical cross section, most beds show a well-developed pebble imbrication. The *a*-axes of the pebbles dip upstream (rather than lying horizontal, transverse to flow) and indicate flow toward the southwest, parallel with the present tectonic strike. This type of fabric is very rare in deposits where pebbles rolled as bed load, because such movement develops a transverse *a*-axis fabric. Bed load transport was minimal to nonexistent for these conglomerates. Alternative mechanisms include some form of mass flow for each bed, or movement of pebbles in suspension. Mass flow is unlikely because of the well-developed imbrication and stratification, and hence it is suggested that the pebbles fell onto the bed out of fluid suspension, and were not subsequently moved. The outstanding problem is how the boulders were maintained in fluid suspension.

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CAMBRIAN OF THE GRAND CANYON—A REEVALUATION

The Grand Canyon Cambrian, previously thought to represent a subtidal transgressive-regressive sequence recording deepening offshore accumulations of marine sandstone, shale, and limestone, is reinterpreted to record shallow-marine, tidal-flat, and fluvial sedimentation on the landward part of a vast cratonic platform marginal to the Cordilleran miogeosyncline.

The basal Tapeats Sandstone is dominantly trough-crossbedded, contains no record of organic activity or marine-tracer grains, and displays a low-variance, unimodal paleocurrent trend down the paleoslope. This part of the Tapeats Sandstone records prevegetation, bed-load fluvial sedimentation.

Shallow-marine "lagoonal" deposits dominate the rest of the sequence: burrowed, very thinly interbedded fine sandstones and shales (Bright Angel Shale) and arenaceous or soft-pellet limestones and dolomitic siltstones (Muav Limestone).

Within the Bright Angel Shale, many 1-6 m thick units of burrowed, cross-laminated sandstone and glauconitic sandstone record shoaling sedimentation. These sandstones locally are succeeded by hematitic oolite beds which were exposure surfaces. Unburrowed, channelled, flaser-bedded sandstones form an extensive 20 m thick tidal-flat sequence in the western Grand Canyon.

Within the Muav Limestone, 1-8 m thick units of dolomitized eocrinoidal biocalcareous and algal-ball limestone, flat-pebble intraclast beds, and a few stromatolites emphasize the shoaling nature of the carbonate platform. A 20 m thick, dolo-