

include waters from the Lake Gosiute basin, late in the history of Lake Gosiute, led to a significant increase in water supplies to Lake Uinta. Initially, the sediment load of this water was trapped in the greater Green River basin during the final stages of Lake Gosiute, so that the water from the north was devoid of clastic debris. The influx of this water, nearly devoid of terrigenous material, resulted in high biologic productivity and the deposition of rich oil shales (Mahogany Bed) in the Piceance basin.

Facies patterns in the Green River Formation in the Piceance Creek and Uinta basins suggest that these two basins were separate lakes until the addition of water from the Lake Gosiute hydrographic basin. Prior to the merging of the lakes in the Piceance and Uinta basins, the brine evolution and, hence, the saline mineralogy were different in the two basins. In the Piceance Creek basin the evaporite minerals were sodium carbonates and chlorides, whereas, in the Uinta basin, the evaporite minerals were sulfate-rich. After merging of the lakes into greater Lake Uinta, the evaporite facies were characterized by sodium carbonate minerals.

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Biostratigraphy of Early and Earliest Late Cretaceous Ostracoda from Peninsular Florida

Deep wells in central and southern Florida have yielded 68 species of marine and brackish-water Ostracoda of Early Cretaceous and earliest Late Cretaceous ages. About 40% of the ostracods are referable to described species. The rock sequence is more than 2,250 m thick.

A partly oolitic limestone facies of the Washitan Stage (early Cenomanian-late Albian) contains 17 species, of which 15 are restricted to the unit. Brackish-water, and perhaps freshwater, as well as marine Ostracoda are represented.

Pre-Washitan Cretaceous rocks of peninsular Florida are principally massive, interbedded carbonate rocks and evaporites and thin shales. Ostracods occur chiefly in the shales. Of 23 species in the Fredericksburgian Stage (middle Albian), nine are restricted to the unit; marine and a few brackish-water species are represented.

Trinitian Stage (early Albian-late Aptian) ostracods are represented by 21 species, of which seven are restricted to that unit; several are brackish-water forms.

Coahuilan Series (early Aptian-Neocomian) rocks contain 35 species, of which 25 are confined to that unit; several are brackish-water types.

The environment represented by the ostracod populations is mainly that of an open-marine shelf bordered by partly brackish-water lagoons. Trinitian Stage rocks contain representatives of a few species which indicate outer-shelf or slope environments.

The population as a whole has strongest affinities for Cretaceous species of the Gulf coastal or Atlantic coastal United States. Several of the Coahuilan and Trinitian species show relationships to European and to South American forms. Few such relations are shown by Fredericksburgian and Washitan species.

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Normally and Reversely Graded Beds Within Large-Scale Foresets of Oolitic Lake-Margin Bench Sequence—Shoofly Oolite of Southwestern Idaho

Pliocene lacustrine oolites composing the lower part of the Glens Ferry Formation crop out in a 40-km northwest-southeast-trending belt along the southwestern margin of the Snake River plain. Near Oreana, Idaho, where this limestone reaches 40 m in thickness, the oolite occurs as three progradational sequences, each consisting of thinly coated ooids in foreset beds up to 20 m thick; these beds are abruptly overlain by thickly coated, massive, burrowed oolite. Foreset beds, each 5 to 15 cm thick, dip basinward to the northeast at 30°; they exhibit both coarsening- and fining-upward trends. Reversed grading (coarsening upward) occurs high in each foreset unit, but the beds become normally graded (fining upward) near the base.

The Shoofly Oolite was deposited as three progradational bench sequences which built lakeward during short periods of stillstand in a longer transgressive phase of Lake Idaho. As such, each bench sequence is analogous to a "Gilbert" delta which extended laterally along the lake margin, but was fed by littoral sands which became coated during transport on the bench platform. Deposition of foreset beds by grain flow on the upper parts of the bench slope, and by fluidized sediment flow on the lower parts of the bench slope, resulted in the formation of reversely graded beds near the bench platform and normally graded beds near the base of the slope. During periods of rising lake level, but prior to the deposition of a subsequent bench sequence, abandoned bench platforms were extensively burrowed and winnowed by waves. As a result, deposition of cortical laminae on platform ooids continued, and the massive oolites which now cap each progradational bench sequence were formed.

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South Hallettsville; Gas Field in Lower Wilcox, Lavaca County, Texas

South Hallettsville field, discovered in 1976, produces gas and condensate from a lower Wilcox clastic wedge. The discovery well, General Crude Oil 1 Anderson, was drilled on an Edwards seismic closure, encountered high pressure gas sands in the lower Wilcox, blew out, and was abandoned. The replacement well had potential of 23,926 MCFGD flowing and 38.7 bbl condensate per MMCF of gas through perforations from 9,909 to 10,011 ft (2,973 to 3,003 m). Within 2 years this well had produced 2.1 Bcf of gas and 74,543 bbl of condensate.

The wedge has a sandstone shale ratio of 20%. Reservoir sands are medium to very fine grained, individual or stacked, and several inches to 25 ft (7.5 m) thick. The sand is 40 to 70% quartz with 10 to 35% feldspar and lithic fragments. The matrix is kaolinite, chlorite, illite, and illite/smectite with minor carbonate cement. Reservoir sands have porosities of 18 to 24.7% and permeabil-