

to be straight or have a low sinuosity with sediment fills ranging from fine mudstones to coarse-grained sandstones. Delta-plain deposits consist of lagoonal, bayfill mudstones and small-splay or bayhead-delta sandstones formed in areas behind and marginal to the delta front. Numerous coarsening-upward sequences are capped by localized coals. Fluvial and upper delta-plain areas consist of channel, levee, and backswamp materials that are laterally discontinuous.

Features such as those observed in these wave-dominated delta deposits are easily recognizable on seismic lines. The seismic lines can be used to target favorable areas of hydrocarbon and coal accumulation.

LEVIN, DAVID M., Gulf Energy and Minerals Company—U.S., Houston, TX

#### Hydrocarbon Exploration in Western Approaches, Offshore England

The North Sea has matured into its production phase and explorationists are now searching for hydrocarbons on the Atlantic continental margin west of England. The Western Approaches, one of these new exploration frontiers, is currently the subject of drilling to test hydrocarbon potential.

The Western Approaches forms an ENE-WSW trending structural trough southwest of England extending from the mouth of the English Channel westward to the edge of the continental shelf. The basin is believed to be the failed arm of a triple junction which originated in Permian-Triassic rifting associated with separation of the North American continent from western Europe and the opening of the Atlantic.

Seismic and gravity data indicate good sediment thickness in which Permian-Triassic, Jurassic, Cretaceous, and Tertiary sequences have been interpreted. Structure of the basin has been strongly influenced by Hercynian related tectonism in basement rocks. Four wells have been drilled within the last 18 months as the first exploratory attempts in this basin.

LINK, MARTIN H., Los Angeles Harbor College, Wilmington, CA, and JOANN E. WELTON, Chevron Oil Field Research Co., Brea, CA

#### Hydrocarbon Potential of Matilija Sandstone, an Eocene Sand-Rich, Deep-Sea Fan and Shallow-Marine Complex, California

The Matilija Sandstone Member, exposed in the Santa Ynez Mountains, California, records a major regressive event in the Eocene Santa Ynez basin in which turbidites were deposited in the basin and subsequently covered by shallow-marine complexes. Despite thick favorable source beds and generally good initial reservoir characteristics, the Matilija sandstone is not a productive unit in the basin. Lowered reservoir rock permeability (<1 md) and porosity (0-10%) are due to early compaction, cementation, and diagenesis.

The lower part of the Matilija sandstone is a 700-m-thick sand-rich deep-sea fan complex which overlies basin plain and turbidite deposits (Juncal shale). The Matilija sandstone consists of anastomosing outer-fan depositional lobes overlain by channelized middle- and

inner-fan deposits. Cross-bedded sandstone, red-bed, and carbonate-evaporite sequences overlie the turbidites. Matilija sandstone deposition closed with rapid transgression which culminated in the deposition of basin plain and turbidite deposits (Cozy Dell shale).

The Matilija sandstone lower deep-sea fan complex has a high sandstone and shale ratio (4:1) and consists of submature arkoses of facies B. The average sandstone is medium grained, moderately sorted, subangular, massive, and contains 40% quartz, 35% feldspar (about equal amounts of potassium and plagioclase feldspars), 10% lithic fragments (mostly granitic and volcanic types), and smaller amounts of mica, chert, and heavy minerals. Early compaction, carbonate cementation, and authigenic pore-lining chlorite, albite, and quartz have reduced the initial porosity and permeability. Minor secondary fractures are the only effective porosity in these rocks now.

LORSONG, J. A., Saskatchewan Geol. Survey, Regina, Sask.

#### Nearshore Lithofacies of Mannville Group, Lloydminster Heavy Oil Area, Saskatchewan

The Lower Cretaceous Mannville Group comprises a 200-m succession of poorly consolidated, fine-grained sandstones and shales in the Lloydminster heavy oil producing area of west-central Saskatchewan. Detailed study of closely spaced cores from several oil fields indicates the presence of six major lithofacies (here denoted by letters), and suggests some provisional interpretations of depositional environments.

Facies L consists of well-sorted sandstone characterized by low-angle cross-lamination and hummocky lamination. Nearly all oil production is from multistory facies L sandstone bodies, which average 5 m in thickness and commonly pinch out over a few hundred meters. Facies T includes moderately sorted sandstones with multidirectional trough cross-lamination. Facies M is composed of massive, fine to medium-grained, poorly sorted sandstone. Facies B comprises bioturbated sandstone-shale sequences with abundant oscillation ripples, desiccation cracks, and flaser, lenticular, and wavy bedding. Facies S includes two subfacies: massive shales (S<sub>1</sub>) are commonly associated with laminated shales (S<sub>2</sub>) that display desiccation cracks as well as flaser, pin-stripe, and tidal bedding. Facies C comprises thin lignite beds composed of terrestrial plant debris.

Sedimentary structures in facies B and S indicate shallow-water deposition with intermittent exposure typical of tidal flats. Characteristic structures of facies L and T suggest beach, offshore bar, or sand-flat depositional environments. Facies M may represent thin channel fills. Intimate association of the lithofacies in vertical section indicates a nearshore depositional setting for the Mannville Group. A substantial number of petroleum reservoirs in the Lloydminster area appear to be intertidal sand bodies rather than the fluvial channel fills suggested in some previous studies.

LORSONG, J. A., Saskatchewan Geol. Survey, Regina, Sask.

### Synsedimentary Deformation in Fossil Accretionary Prism, Greece

Structural and stratigraphic evidence demonstrates that Paleogene turbidites of the Polistes Formation were progressively deformed during sedimentation. The Polistes Formation is preserved in internally undeformed thrust sheets that are tectonically intercalated with sheets of severely deformed blocks of the formation. This imbricate stack is interpreted as a fossil accretionary prism.

The Polistes Formation typically consists of hemipelagic red shales and limestones (basin-plain facies association) followed by thin-bedded terrigenous turbidites (fan fringe, interlobe and distal depositional lobe facies associations) which are overlain then by predominantly thick-bedded turbidites (depositional lobe facies association). This "normal" progradational sequence is interrupted by incised channel complexes which lie above basin-plain sediments and beneath fan-fringe deposits in some thrust sheets. Atypical facies organization suggests tectonic activity during sedimentation. Small-scale soft sediment deformation in the form of convolute lamination exists in about 20% of turbidite sandstone beds. Assuming that the deformation resulted from seismic activity, sedimentation rates and the distribution of structures indicate that seismic shocks affected the depositional area every 10 to 100 years on the average. The most important evidence for synsedimentary deformation lies in the distribution of marl marker beds among thrust sheets. The stratigraphic distribution of marls corresponds uniquely to individual thrust sheets. This and other sedimentological relations demonstrate that marl and turbidite accumulation was controlled by progressive tectonic removal of thrust sheets from active deposition. The absence of unconformities within thrust sheets suggests that the deep sea fans represented in the Polistes Formation were deposited in a trench.

### LOUCKS, R. G., and M. M. DODGE, Bur. Econ. Geology, Austin, TX

#### Sandstone Diagenesis in Geopressured Tertiary Gulf Coast Basin

Texas Gulf Coast sandstones have been subjected to several stages of burial diagenesis: mechanical compaction, silica and carbonate cementation, and cement and grain dissolution. Each stage is well-developed before the sandstone aquifer is geopressured; only iron-rich carbonate cement continues to be precipitated in the geopressured zone.

Fluids must pass through an aquifer for cementation to occur. Ideally there is minimal fluid movement in the geopressured zone because overpressuring requires the fluids to be trapped. However, if faults are periodically opened and fluids escape, a drop in pressure of the aquifer would result. Lower pressure would allow carbon dioxide gas to evolve and the carbonate equilibrium would then favor precipitation of carbonate minerals.

Evidence of the above process is scale precipitated in wells producing geopressured fluids. As the fluid pressure is lowered by production of the well, iron-rich car-

bonate minerals are formed. These carbonate minerals are precipitated because of a pressure drop which may be analogous to pressure drops along opened fault zones.

### LOW, PHILIP F., Purdue Univ., West Lafayette, IN

#### Properties of Water in Clay Mineral Systems

Among the properties of water in clay mineral systems receiving special attention are specific volume, specific heat capacity, coefficient of thermal expansion, and viscosity. Every property,  $J$ , of the interstitial water is described by the equation  $J = J^{\circ} \exp [\beta/m_w/m_m]$  in which  $J^{\circ}$  is the value of the property for pure bulk water,  $\beta$  is a constant and  $m_w/m_m$  is the mass ratio of water to montmorillonite—the clay mineral used as a prototype.

The swelling of clay will also be discussed and it will be shown that  $m_w/m_m = \lambda \rho_w S / 2(1-r)$  where  $m_w/m_m$  is the mass ratio of water to montmorillonite (or any other layer silicate),  $\lambda$  is the interlayer distance,  $\rho_w$  is the density of the interlayer water,  $S$  is the specific surface area of the clay and  $(1-r)$  is the fraction of the water in interlayer regions. In this equation,  $\lambda$  is a function of the swelling pressure,  $\pi$ . If  $\lambda$  is the same for all clays at any given  $\pi$ , the equation indicates that a plot of  $m_w/m_m$  against  $S$  for different clays yields a straight line. Experimental data show that this is the case. Then the value of  $\lambda$  at each of several values of  $\pi$  can be calculated from the slope,  $\lambda \rho_w / 2(1-r)$ , of the corresponding straight line and the results used to make plot of  $\pi$  versus  $\lambda$  which should be valid for all clays. Finally, this plot and the foregoing equation are used to show how excess pressure develops in a shale from which water cannot escape when the geostatic pressure increases or when  $S$  decreases owing to the conversion of montmorillonite to illite.

### LOWRIE, ALLEN, and WILLIAM P. SEARCY, III, U.S. Naval Oceanog. Office, Bay St. Louis, MS

#### Multi-Parameter Regional Mapping and Interpretations

A global mapping program of those geologic parameters that may influence acoustic propagation through the ocean floor has evolved a number of mappable parameters including bathymetry, physiographic provinces, surface sediments, sediment thickness, paleomagnetic anomalies and age of seafloor, inferred thickness of lithosphere, heat flow, potential temperature of deepest waters, areas of possible bottom current activity, acoustic velocity of deepest waters, acoustic velocity in uppermost sediments, calculated ratios of acoustic velocity along water-sediments interfaces, possible petrologic provinces within lithosphere, roughness of acoustic basement, regional structure, and earthquake locations.

Mapping efforts present maximum detail consistent with data and application of established geologic principles. In areas lacking sufficiently definitive data, a reasonable interpretive fabric is shown. By interrelating appropriate parameters, greater interpretive accuracy is routinely possible.

The principal correlations between parameters are