cally regardless of the price of crude oil. One exception to this economic situation is shale oil coproduced with other minerals from the vast "Saline Zone" oil shale resource of northwest Colorado.

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Wattenberg Field, Denver Basin, Colorado

The most important mineral resource activity in Colorado during the past decade has been the discovery and development of the Wattenberg gas field. Located north of Denver near the axis of the Denver basin, Wattenberg is estimated to have reserves of 1.3 tcf in the tight J Sandstone reservoir over an area of 600,000 acres  $(2,400 \, \mathrm{km}^2)$ , at depths between 7,600 and 8,400 ft (2,310 and 2,560 m). Net pay thickness varies from 10 to 50 ft (3 to 15 m), porosity ranges from 8 to 12%, and permeability varies from 0.05 to 0.0005 md

The J is interpreted as a fluvial-deltaic sandstone with the principal production from widespread delta-front sandstone. Drilling for gas in the Cretaceous J Sandstone has resulted in multiple pays in overlying strata. The Spindle field, situated in the southwest portion of the Wattenberg field, produces from two marine bar complexes (Hygiene and Terry) in the middle portion of the Pierre Shale. Since 1971, total production is in excess of 36 million bbl of oil and 164 bcf of gas from depths of 4,000 to 5,000 ft (1,220 to 1,525 m).

For the past 2 years, the Codell Sandstone, approximately 500 ft (152 m) stratigraphically above the J Sandstone, has been developed as a new petroleum-producing zone. More than 100 discoveries have been made within and marginal to the outlined Wattenberg field area. The Codell is a tight bioturbated marine-shelf sandstone generally without a central-bar facies. Net pay thickness ranges between 3 and 25 ft (0.9 and 7.6 m), porosities range between 8 and 24% (with the average 10–12%), and permeabilities less than 0.5 md. Because of rapid decline in production and economic uncertainties, potential reserves from the Codell are unknown.

All petroleum accumulations in the Wattenberg area are regarded as stratigraphic traps. Variation in pay thickness and reservoir quality is related to the original environments of deposition and paleostructure which locally influenced unconformities, fracturing and diagenesis.

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Shallowing-Upward Sequence in a Pleistocene Coral Reef and Associated Facies, San Salvador, Bahamas

Excellent exposures of a well-preserved Pleistocene coral reef (dated 120 to 140 thousand years B.P.) extend along the shore northwest of Cockburn Town on the island of San Salvador, which lies at the eastern edge of the Bahamas platform 600 km (375 mi) east-southeast of Miami, Florida. The reef was buried sequentially by nearshore marine and terrestrial carbonate sands during a lowering of sea level. This shoaling produced a vertical sequence of carbonate rocks that represents a progressive change from shallow subtidal through beach to dunel environments.

Acropora palmata, A. cervicornis, Diploria sp., and Montastrea annularis are found in growth position up to 2 m (6 ft) above present mean sea level (MSL). Coral rubble, composed principally of A. cervicornis and A. palmata, commonly occurs subjacent to, adjacent to, and up to 80 cm (31 in.) above the in-situ corals. Gullies in the coral rudites are filled by coarse skeletal calcarenites, which in some places contain well-preserved, irregular boxworks of Ophiomorpha sp. and Skolithos linearis burrows. The calcarenites extend upward and outward from the gullies to overlie the corals and coral rudites; lowermost beds commonly have trough cross bedding extending up to 3.6 m (12 ft) above MSL. Upward, the calcarenites become finer grained, skeletal, intraclastic, peloidal, and oolitic, with gently seaward dipping planar cross beds and beachrock clast breccias occurring up to 4 m (13 ft) above MSL. Fine-grained eolianites with abundant rhizocretions and some paleosols occur higher than 4 m (13 ft) above MSL, and this elevation marks the maximum height of sea level during the development of the Cockburn Town reef. Following this high stand, lowering of sea level produced this fine Pleistocene example of a shallowing-upward carbonate sequence.

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Fluvioglacial Sandstone Reservoirs and Deposystem Analysis in Hydrocarbon Exploration of Permian Gidgealpa Group, Southern Cooper Basin, South Australia

The sedimentology of the Permian Gidgealpa Group of the southern Cooper basin currently is being evaluated to ascertain the tectonosedimentologic evolution of the basin and to determine the architecture of the clastic suite in order to generate exploration plays.

The study has produced a new understanding of the relationship between the hydrocarbon-rich Tirrawarra Sandstone, of dominantly fluvial origin, and the, as yet little explored, glacigenic Merrimelia Formation. The facies states and transitions of both formations interdigitate, and typical porous Tirrawara-type fluvial facies are clearly evident within glaciolacustrine Merrimelia-type sediments.

The Merrimelia Formation was examined regionally in 29 cored wells. The formation attains a maximum thickness of 300 m (1,000 ft), and representative facies include glaciofluvial outwash, terrestrial and subaqueous diamictites, and glaciolacustrine, wave-affected, and ripple-laminated sandstones, with thick, monotonous mudrock sequences containing clay-dominant rhythmite horizons.

The Tirrawarra Sandstone, analyzed in 32 cored wells, comprises four major facies associations throughout its maximum 75 m (250 ft) thickness. These associations indicate a temporal and spatial evolution of a fluvioglacial to predominantly fluvial system. Initial deposition on low slope, outwash fans, where braided processes operated is indicated. This sedimentation style evolved into a low sinuosity, bedload-dominant, sandy braided system, with high width-to-depth ratio channels. Allocyclic control mechanisms are invoked for "late Tirrawarra" sedimentation as the facies reveal proximal-distal patterns and the fluvial style changes to a mixed-load channel system.

The interfacing and evolutionary pattern of the deposystem indicates that additional reserves potential exists for reservoirs developed locally within the Merrimelia Formation.

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Salt Domes, Anoxic Brines, and Organic-Rich Sediments in Intraslope Basins

The origin of organic-rich sediments and their potential relationship to petroleum source beds remains problematic despite the development of models relying primarily on oceanographic mechanisms. We outline instead a model that involves salt dome development under a prograding shelf-slope and has a modern analog in the hypersaline anoxic Orca basin on the Texas-Louisiana continental slope. We believe the model may explain the occurrence of some ancient black shale sequences by a mechanism which is not dependent on either intensification of the mid-water oxygen minimum, enhanced marine productivity, or circulation-induced bottom water anoxia. In our working model, salt diapir movement structurally produces sedimentary basins in the intraslope region. Dissolution of the salt inhibits oxygen replenishment in the deep parts of the basin. As oxygen levels and bioturbation in the basin are diminished, finely laminated sediments with high organic carbon contents result. Gradual evolution of the basin to complete anoxic conditions leads to the accumulation of fine-grained, fluid-rich, black muds which, upon burial and dewatering through compaction, resemble black anoxic shales. Infilling and burial of the basin with clastic and hemipelagic sediments from glacioeustatic-driven climatic episodes subsequently provide source beds and reservoir rocks for potential exploration plays. Lerche and O'Brien have recently shown that the high thermal heat capacity of salt domes (≈5 times > sediment) produce predictable thermal anomalies in the surrounding sediments. Thermal maturation rates of the organic-rich source beds are enhanced near the top of the salt dome and suppressed on the lower flanks. The negative thermal anomaly on the lower flanks of the salt dome inhibits overmaturation and thus enlarges the hydrocarbon window. As overburden increases, the density contrast between the salt ( $\approx 2.2 \,\mathrm{gm/cm^3}$ ) and surrounding compacted sediments (2.6-2.7 gm/cm<sup>3</sup>) eventually becomes such that upward diapirism resumes. The present model, although preliminary, may have wide applicability in sedimentary basins other than the Gulf of Mexico. Further work on the Orca basin and other sedimentary sections near salt displacement should prove beneficial in future modeling of organic-rich sediments.